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# Miocene fire intensification linked to continuous aridification on the Tibetan Plateau

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## ABSTRACT

Although fire is considered an important factor in global vegetation evolution and climate change, few high-resolution Miocene fire records have been obtained worldwide. Here, two independent micro-charcoal-based fire records from the northern Tibetan Plateau were analyzed; both show similar trends in micro-charcoal concentrations through time, with low abundances in the warmer Middle Miocene Climate Optimum (18–14 Ma) followed by a continuous increase throughout the late Miocene (14–5 Ma) cooling. Our detailed statistical analyses show that the micro-charcoal concentration trend is highly positively correlated to the trend in oxygen isotopes ( $\delta^{18}\text{O}$ ,  $r = 0.94$ ) and xerophytic species ( $\%_{\text{xero}}$ ,  $r = 0.95$ ). We propose that the intensified fire frequency on the Tibetan Plateau mainly originated from the forest-steppe ecotone as a result of the continuous aridification in winter driven by the global cooling and decreased atmospheric  $p\text{CO}_2$  that occurred during 18–5 Ma, with a secondary control by the tectonic activity of the northern Tibetan Plateau.

## INTRODUCTION

Late Cenozoic global fire activity has been considered as one of the key factors influencing modern vegetation evolution and climate change (Bowman et al., 2009; Dale et al., 2001;

Edwards et al., 2010). However, globally, few Miocene high-resolution records have been obtained with which to assess fire occurrence (Hoetzel et al., 2013). As part of the second-largest arid to semi-arid area in the world, the northern Tibetan Plateau is a unique location for studying fire history, along with vegetation and aridity evolution (Miao et al., 2016); however, the region's numbers of fire records, sampling resolution, combustion source data, and controlling factors still need further investigation.

The charcoals resulting when plant matter is burned have been referred to as “fossils of fire” and are regarded as one of the most important proxies in the study of fire history (e.g., intensity, frequency, etc.) (Patterson et al., 1987; Whitlock and Larsen, 2002; Bowman et al., 2009; Bond, 2015). Here, we analyze fire histories based on relatively high-resolution micro-charcoal records from two independent sites on the northern Tibetan Plateau: the KC-1 core (38°03'N, 91°45'E; drilled by the Qinghai Petroleum Administration Bureau) in the Qaidam Basin, characteristic of a lacustrine environment (Miao et al., 2011), and the KM/KNX sections in the Kumkol Basin, characteristic of a fluvial-lacustrine environment (Lu et al., 2016, 2018) (Fig. 1).

## STUDY SITES AND DATA

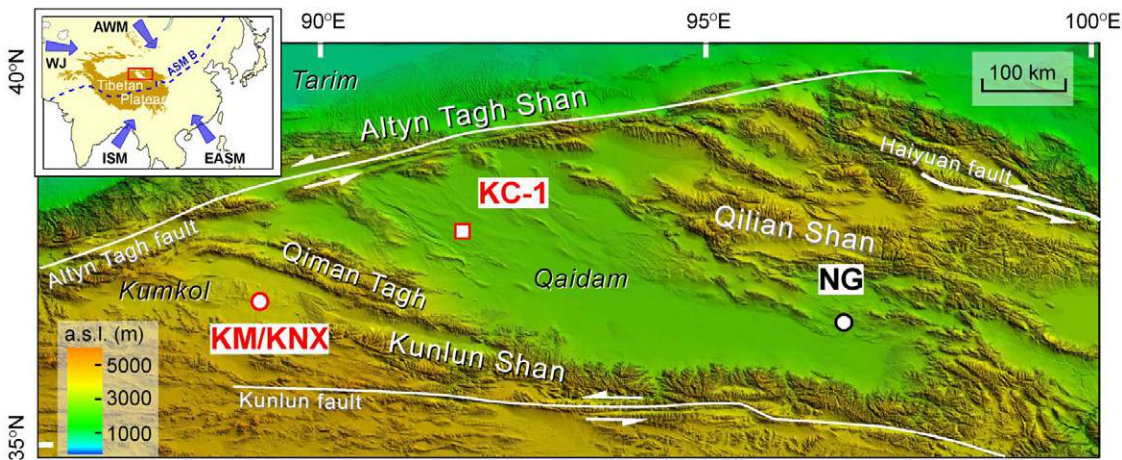
Situated in southern Inner Asia, the northern Tibetan Plateau is marked by a typical inland hyperarid climate, with mean annual

precipitation of <200 mm. As a result of its high elevation, the area is also very cold and windy. The Westerlies represent the dominant atmospheric circulation pattern throughout the year, while the East Asian summer monsoon generally reaches only the southeastern Qaidam Basin (Wu et al., 1985) (Fig. 1). Under these conditions, bare ridges and desert vegetation constitute the typical land cover. The desert vegetation that grows on gravelly lake margins and on well-drained soils of alluvial fans is composed mainly of *Ephedra*, *Nitraria*, *Tamarix*, Poaceae, Cyperaceae (as well as Asteraceae), and occasionally *Picea* and *Sabina* on the high mountains near the eastern part of the Qaidam Basin.

Tectonically, the ~1600-km-long Altyn Tagh fault separates the low-lying Tarim Basin to the north from the Tibetan Plateau. The left-lateral strike-slip Kunlun fault marks the northern boundary of the high-elevation, low-relief main part of the Tibetan Plateau, extending for ~1500 km. A broad triangular area consisting of the Kumkol Basin, the Qiman Tagh range, and the Qaidam Basin is bounded to the northwest by the Altyn Tagh fault and to the south by the Kunlun fault (Fig. 1). Previous work based on apatite fission-track modeling has shown that the Qaidam and Kumkol Basins have been separated by the synchronous initial uplift of Qiman Tagh at ca. 40–30 Ma (Liu et al., 2017).

The present Qaidam Basin, with an average elevation of ~3000 m, is the largest intermontane basin on the northeastern Tibetan Plateau, and

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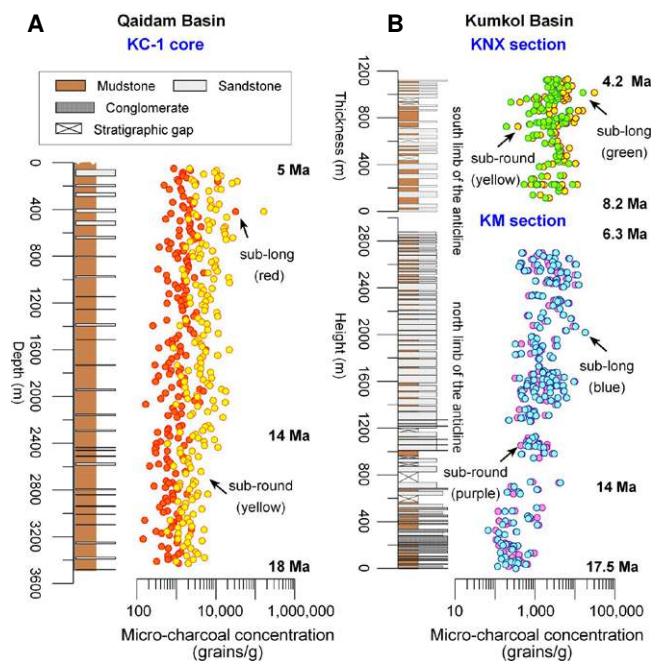
**Figure 1.** Location map showing northern Tibetan Plateau, major strike-slip faults, and sites of micro-charcoal study. KM/KNX sections and KC-1 core are located in the Kumkol and western Qaidam Basins, respectively. NG section lies in the eastern Qaidam Basin (Miao et al., 2016). AWM—Asian winter monsoon; ASM B—Asian summer monsoon boundary; EASM—East Asian summer monsoon; ISM—Indian summer monsoon; WJ—Westerly jet.

covers an area of  $\sim 120,000$  km<sup>2</sup>. The sediments recovered from the KC-1 core, reaching 3425 m in depth, were extracted from the western basin and span ages of 18 Ma to 5 Ma, with an average sedimentation rate of  $\sim 260$  m/m.y. (Miao et al., 2011) (Fig. 2A). Another section, named NG (Naoge), from the eastern basin and spanning ages of ca. 18–10 Ma, was also used for comparison (for details, see Miao et al., 2016). The rhombus-shaped Kumkol Basin has an average elevation of  $\sim 4200$  m and an area of  $\sim 17,500$  km<sup>2</sup> (Fig. 1). We sampled the KM section, spanning 2869 m in thickness, and the KNX section, spanning 1125 m, from the central basin. Chronostratigraphic analysis of these two sections yielded ages of 17.5–6.3 Ma in the KM section (Lu et al., 2016), and 8.2–4.2 Ma in the KNX section (Lu et al., 2018), providing average sedimentation rates of  $\sim 260$  m/m.y. and  $\sim 290$  m/m.y., respectively (Fig. 2B). The composite age range of 17.5–5.0 Ma for these two sections (herein KM/KNX) was used for the micro-charcoal analyses and comparison with the KC-1 core.

## METHODS AND RESULTS

In this study, 188 samples from the KC-1 core (Qaidam Basin) and 250 samples from the KM/KNX sections (Kumkol Basin) were used for micro-charcoal identification and processed via standard palynological techniques. A known number of *Lycopodium clavatum* spores were initially added to each sample for calculating the micro-charcoal concentration (MC) (Miao et al., 2016). Two basic micro-charcoal shapes (sub-long [*L*] and sub-round [*R*]) were identified by calculating the ratio of length (major axis) to width (minor axis): if the value was  $>2.5$ , the micro-charcoal grain was classified as *L*, and if  $<2.5$ , as *R*. The MCs for both types ( $MC_L$  and  $MC_R$ , respectively) were determined, and then added together to obtain the total MC values, referred to as  $MC_{total}$  (Table DR4 in the GSA Data Repository<sup>1</sup>). The  $MC_{total}$  for the previously

**Figure 2. Micro-charcoal records and lithology (see Fig. 1 for location).** A: KC-1 core, western Qaidam Basin, northern Tibetan Plateau, dominated by lacustrine sediments. B: Composite KM and KNX sections, Kumkol Basin, dominated by fluvial-lacustrine sediments. Colored dots indicate micro-charcoal concentration (MC), in grains per gram; subscripts *L* and *R* refer to sub-long and sub-round grains, respectively.



studied NG section in the eastern Qaidam Basin (Miao et al., 2016; Fig. 1) was also included in the statistical analyses.

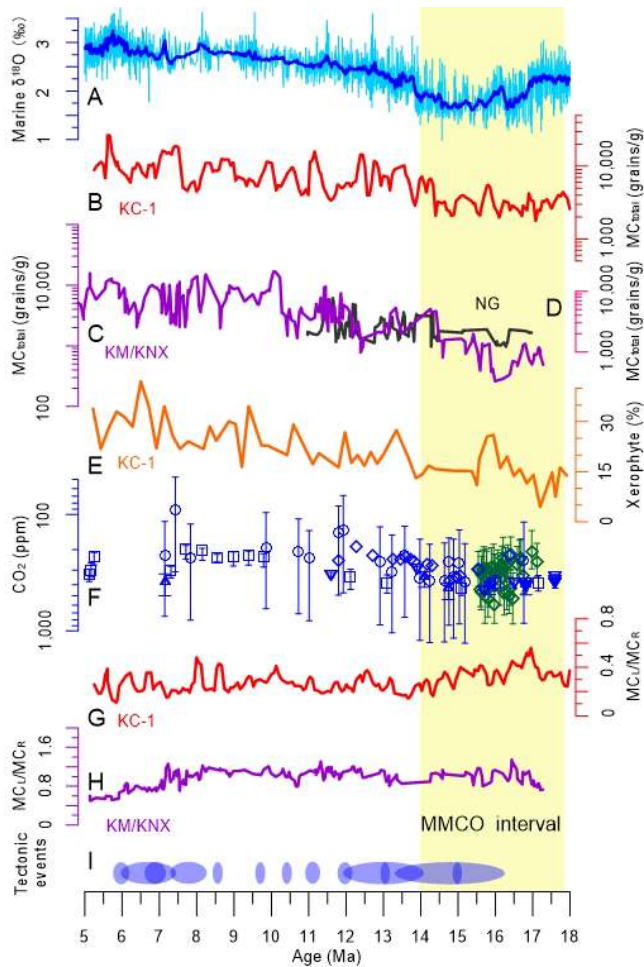
First, the existence of a trend in the three separate time series (KC-1, KM/KNX, and NG) of observed  $MC_{total}$  values was tested using the Mann-Kendall trend test (Table DR1). Then Pearson correlations were computed for the three possible pairs of  $MC_{total}$  time series based on spline-fitted log-transformed  $MC_{total}$  values interpolated every 50 k.y. along each time series, resulting in three interpolated data sets described within the same time frame (Table DR2). Last, a composite  $MC_{total}$  time series was compiled from the three separate series, and finally correlated to the time series of marine  $\delta^{18}O$ , atmospheric  $pCO_2$ , and percentages of xerophytic taxa in the Qaidam Basin using the same interpolation procedure based on spline-fitted curves (Table

DR3). In all cases, correlation significance was evaluated through a Mantel test based either on the interpolated  $MC_{total}$  values or on the time autocorrelation-free first-order differences between “raw” interpolated  $MC_{total}$  values. All computations were done using PAST software, version 3.16 (<https://folk.uio.no/ohammer/past/>; Hammer et al., 2001).

The most notable characteristics between the KC-1 and KM/KNX sites are their similar patterns (Pearson  $r = 0.95$ ; Table DR2) but different absolute values of  $MC_R$ ,  $MC_L$ , and  $MC_{total}$ . The KC-1 core has the highest grain counts, varying between 46,300 and 900 grains/g, with an average of  $\sim 7000$  grains/g (Fig. 2A). The KM/KNX sections vary between 36,100 and 200 grains/g, with an average of  $\sim 4000$  grains/g (Fig. 2B). Overall, the trends of  $MC_{total}$  in the three sections (including NG) are all highly significantly increasing with time (Figs. 3B–3D; Tables DR1 and DR2). As a result, the composite  $MC_{total}$  time series also shows a highly significant trend (Mann-Kendall

<sup>1</sup>GSA Data Repository item 2019105, micro-charcoal data and details of the statistical analyses, is available online at <http://www.geosociety.org/datarepository/2019/>, or on request from [editing@geosociety.org](mailto:editing@geosociety.org).

**Figure 3. Data time series of climatic proxies and tectonic events on the northern Tibetan Plateau. MMCO—Middle Miocene Climate Optimum. A: Global deep-sea  $\delta^{18}\text{O}$  records (Zachos et al., 2001) (light blue) with 12-point averaging (darker blue). B–D: Micro-charcoal concentration (MC) records from KC-1 core, western Qaidam Basin (this study) (B), KM/KNX sections, Kumkol Basin (this study) (C), and NG section, eastern Qaidam Basin (Miao et al., 2016) (D). E: Xerophytic pollen percentages in KC-1 core (Miao et al., 2011). F: Synthesis of published  $p\text{CO}_2$  proxy data. Boron (diamonds) from Foster et al., 2012; Greenop et al., 2014. Stomata (inverted triangles) from Kürschner et al., 2008. Alkenone (regular triangles) from Zhang et al., 2013. B/Ca (squares) from Tripathi et al., 2009. Paleosol (circles) from Ji et al., 2018. Error bars represent uncertainties in underlying assumptions of each proxy. G,H: Ratios of  $\text{MC}_L/\text{MC}_R$  ( $L$ —sub-long grains;  $R$ —sub-round grains) from KC-1 core and KM/KNX sections, respectively (this study). I: Tectonic events on the northern Tibetan Plateau (after Miao et al. [2012], Li et al. [2014], and Chang et al. [2015]). Curves in B–D and G–H are by three-point averaging.**



test:  $S = 54,543$ ;  $Z = 11.88$ ;  $p = 1.5 \times 10^{-32}$ ) corresponding to an increase of approximately one order of magnitude in the spline-fitted  $\text{MC}_{\text{total}}$  curve over the analyzed time interval. This composite spline-fitted  $\text{MC}_{\text{total}}$  curve shows a very high correlation ( $r = 0.94$ ) with the global deep-sea isotope record at a 50 k.y. time resolution (Zachos et al., 2001) (Fig. 3A; Table DR3). Additionally, it also highly correlates ( $r = 0.95$ ) with the xerophytic pollen abundances in the KC-1 core (*Ephedra*, *Nitraria*, *Chenopodiaceae* (Amaranthaceae), *Artemisia* and other Asteraceae taxa, etc.; Miao et al., 2011; Fig. 3E), whereas it is negatively correlated ( $r = -0.86$ ) to the comprehensive  $p\text{CO}_2$  record (Kürschner et al., 2008; Tripathi et al., 2009; Foster et al., 2012; Greenop et al., 2014; Ji et al., 2018; Fig. 3F).

## DISCUSSION

### Fire and Aridification

Based on the relationship between  $\text{MC}_{\text{total}}$  and fire (Hoetzel et al., 2013; Patterson et al., 1987; Bond, 2015), higher values of  $\text{MC}_{\text{total}}$  over long time scales (e.g., millions of years) can be related to increased fire activity, e.g., greater

frequency or intensity. The  $\text{MC}_{\text{total}}$  data from the KC-1 core and KM/KNX sections (Figs. 3B and 3C), as well as from the NG section (Miao et al., 2016) (Fig. 3D), support the view that fire activity on the northern Tibetan Plateau steadily increased during the middle to late Miocene. The overall average expansion of xerophytic vegetation in the western Qaidam Basin (Miao et al., 2011), as well as an increase in  $\text{MC}_{\text{total}}$ , support an increase in fire activity and suggest that these two phenomena are strongly linked from 18 to 5 Ma (Miao et al., 2016). A relatively low fire frequency and/or strength occurred during the relatively wet Middle Miocene Climate Optimum (MMCO, 18–14 Ma), while an overall increasing trend is observed during the drier late Miocene (14–5 Ma).

### Fire Distribution

Although the relationship between fire intensification and aridification on the northern Tibetan Plateau correlates well, the question of fire distribution remains. It is very important to understand whether the fire intensification is due to an enlarged combustion area or an increase of combustion frequencies.

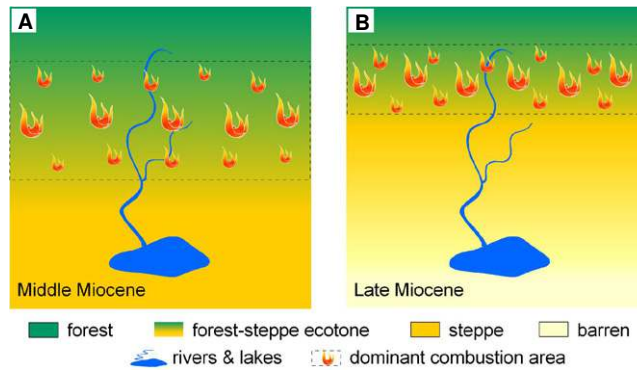
Studies show that the  $L$  and  $R$  in micro-charcoal shapes represent grass and wood sources, respectively (Umbanhowar and Mcgrath, 1998; Daniau et al., 2013; Crawford and Belcher, 2014). This means that the samples with higher  $\text{MC}_L$  seem to be dominated by the combustion of grasses, while the samples with higher  $\text{MC}_R$  seem to be dominated by combustion of wood. Therefore, the ratios of  $\text{MC}_L/\text{MC}_R$  can be directly used to indicate the grass/wood ratios during combustion, with higher values indicating more grass than wood, and vice versa. In this study, the higher values of the  $\text{MC}_L/\text{MC}_R$  in the KM/KNX sections ( $\sim 0.6$ – $1.0$ ) are related to combustion dominated by grass (Fig. 3H), while lower values in the KC-1 core ( $\sim 0$ – $0.5$ ) suggest relatively higher wood combustion (Fig. 3G). Interestingly, although the  $\text{MC}_L/\text{MC}_R$  values are different, the trends remained stable over the examined interval for both sections. Therefore, despite differences in vegetation source material between the two sections, and the expansion of grasslands during late Miocene cooling, the forest-steppe ecotone must have still been the dominant environment where fire occurred in order for such stable trends in  $\text{MC}_L/\text{MC}_R$  values to have been maintained. The conceptual renderings in Figure 4 illustrate this phenomenon. During the MMCO, the area of the forest-steppe ecotone was large but fire frequency was low (Fig. 4A) under the relatively wet climate (Miao et al., 2011), and less micro-charcoals were produced (Fig. 4A). In contrast, during the late Miocene, the forest-steppe ecotone remained a fire source area but likely experienced an overall decreasing trend in areal extent as xerophytic vegetation expanded under a drying climate (Miao et al., 2011). However, the persistently increasing fire frequency due to the drier conditions produced more charcoal particles, thereby greatly counteracting the negative effect of ecotone shrinkage (Fig. 4B).

### Global Cooling, $\text{CO}_2$ , and Tectonics

After compiling global low-sampling-resolution fire records, Bond (2015) argued that none of the potential global factors (oxygen, rainfall seasonality,  $\text{CO}_2$ , novel flammable growth forms) could adequately explain the spread of increased fire activity over long Cenozoic time scales. In this study, the increasing fire frequency is strongly negatively correlated to Bond's compiled results of atmospheric  $p\text{CO}_2$  ( $r = -0.86$ ; Fig. 3F) and is positively closely linked to the global oxygen isotope record (Zachos et al., 2001) ( $r = 0.94$ ; Fig. 3A) and the expansion of xerophytic vegetation ( $r = 0.95$ ; Miao et al., 2011; Fig. 3E).

$\text{CO}_2$  has been argued as having an important role in fertilization of vegetation through affecting photosynthetic assimilation (Wolfe and Erickson, 1993), thus changing the amount of annual gross primary production (Sun et al., 2014). Theoretically, high  $p\text{CO}_2$  facilitates

**Figure 4. Conceptual diagram showing change in fire location on the northern Tibetan Plateau in relation to xerophytic vegetation expansion. A: During warm and wet Middle Miocene, e.g., Middle Miocene Climate Optimum, area experiencing fires was large, but fire frequency was low, causing total micro-charcoal concentration ( $MC_{total}$ ) to be low. B: During late Miocene cooling and drying, area with fire activity shrank but fire frequency became much higher, causing  $MC_{total}$  to steadily increase. In both A and B, fire mainly occurred in forest-steppe ecotone to keep  $MC_L/MC_R$  ( $L$ —sub-long grains;  $R$ —sub-round grains) ratio broadly stable.**



regional vegetation biomass, which, as a result, provides an underlying source for higher MC during combustion, and vice versa. However, the negative correlation between  $CO_2$  and MC may indicate a complicated relationship between  $pCO_2$  and MC.

Global cooling can also be argued as having a major role on the fire through reducing the amount of water vapor held in the atmosphere, favoring drier vegetation (Miao et al., 2012) and, ultimately, increasing combustion frequencies. But the relationship is not always that simple, and seasonal variability has to be taken under consideration. For instance, taking into account the magnetic proxy built up in the Qaidam Basin possibly indicating a wetting event driven by the East Asian summer monsoon at 8.2–7.8 Ma (Nie et al., 2017), we argue that the wet-dry monsoonal climate might have played a major role in promoting fires during the late Miocene. Grasses grow in the wet season and become highly flammable in the dry season; that is, the drier trends most possibly occurred in the winter.

While it is intuitive to link global cooling to increased fire via the role of aridification (Miao et al., 2012), uplift of the Tibetan Plateau has been argued as the primary driver of global cooling and atmospheric  $pCO_2$  decline (Raymo and Ruddiman, 1992; France-Lanord and Derry, 1997; Kump et al., 2000; Galy et al., 2015), which means that fire frequency should have also been clearly increased by the uplift of the Tibetan Plateau, possibly as much as by the atmospheric  $pCO_2$  change. In fact, the uplift of the northeastern Tibetan Plateau is also viewed as a potentially important factor controlling regional aridification (Zhuang et al., 2011; Liu et al., 2015). However, the uplift process is characterized by multiple-stage pulses with enhanced uplift at ca. 15 Ma (Chang et al., 2015), ca. 12 Ma (Wang et al., 2017), and ca. 8 Ma (Fang et al., 2007; Li et al., 2014) (Fig. 3I). The continuous fire intensification has not shown obvious responses to these rapid uplift periods, implying that regional uplift might not be the primary factor controlling fire activity.

## CONCLUSION

A series of detailed micro-charcoal data sets obtained from two new independent sites on the northern Tibetan Plateau (from the KC-1 core in the western Qaidam Basin) and previously published data from the NG section in the eastern Qaidam Basin, offer a perspective on the fire history during 18–5 Ma on the northern Tibetan Plateau. Our results show increased fire occurrence based on low values of micro-charcoal during the warm and less-dry MMCO to high values during the cooler and drier late Miocene. Increased fire frequencies in the forest-steppe ecotone counterbalanced the concurrent reduction in combustion area. Our data and statistical analyses support the view that fire activity on the northern Tibetan Plateau is primarily linked to global cooling, with decreased atmospheric  $pCO_2$  reflecting the continuous aridification on the northern Tibetan Plateau at 18–5 Ma. The influence of tectonic activity on the northern Tibetan Plateau is likely a secondary controlling factor.

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